IDA2017 – A Next-Generation Coupled Modular Assimilation Package

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Outline

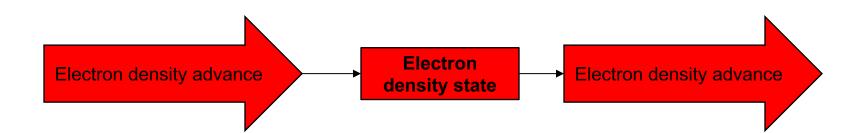
IDA2017 overview and motivation:

- Provide a unified interface for a range of coupled and interlinked geospace models one makefile and one configuration file
- > Automated data download and pre-processing for many data-types
- Develop new tools to integrate diverse observational datasets, ranging from the plasmasphere to the thermosphere

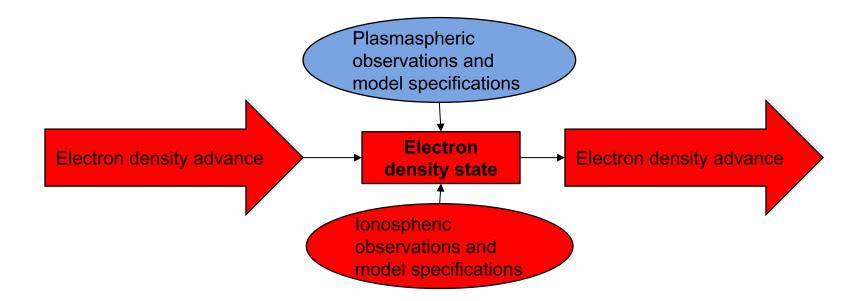
Principal modules:

- Ionospheric physics temporal advance
- > Ionospheric data assimilation
- > Thermospheric composition assimilation
- > Thermospheric wind estimation
- > Plasmaspheric data assimilation

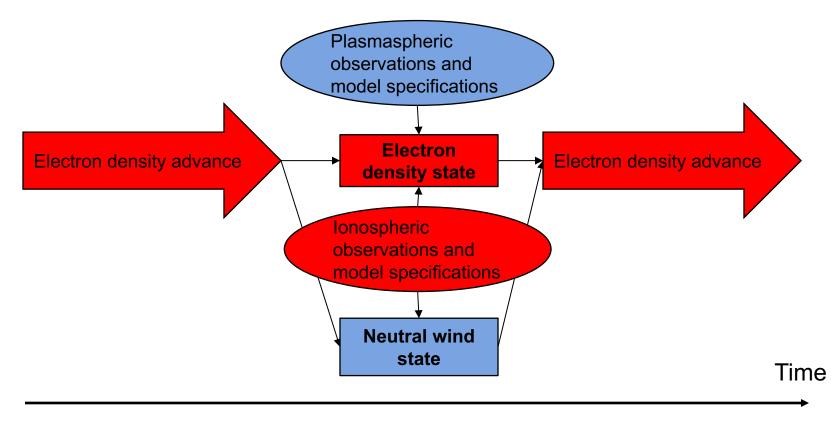


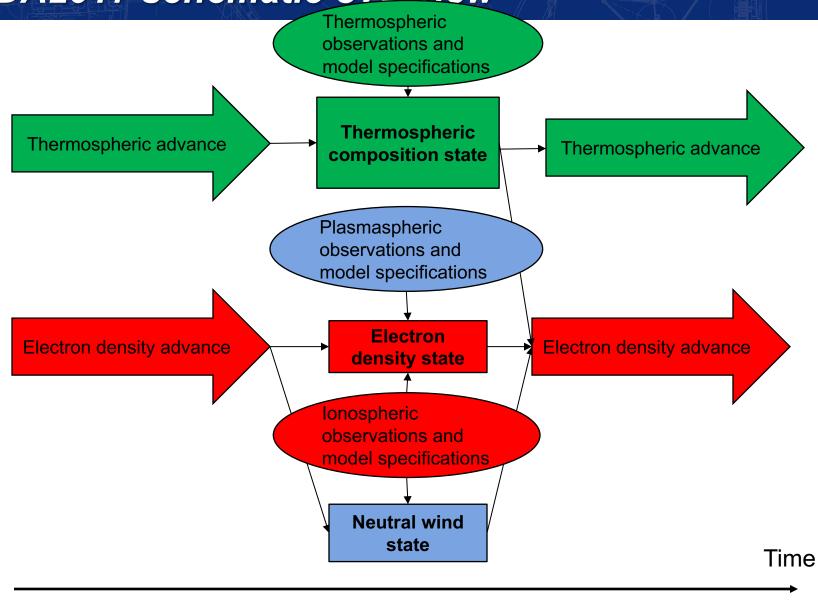




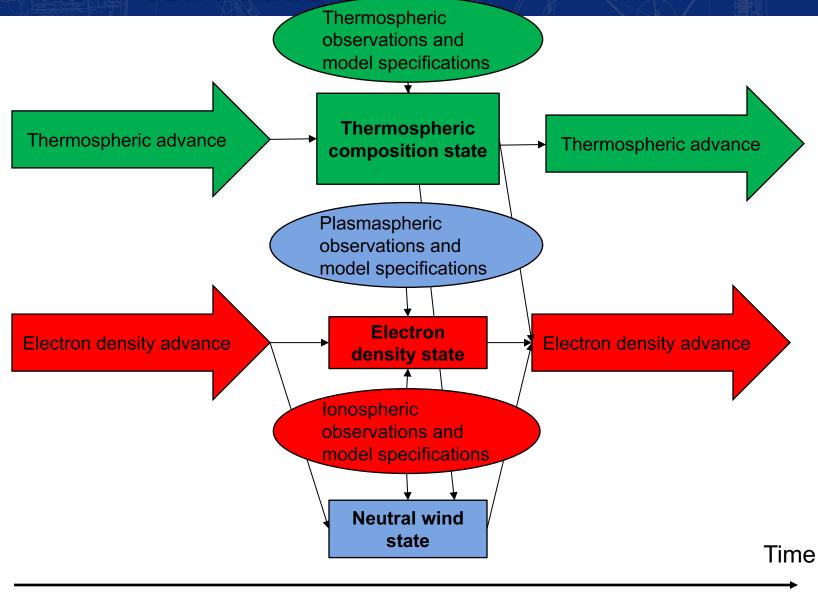


Time





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The ionospheric physics module provides short predictions of ionospheric electron density based on established physics, external drivers and initial conditions

Drivers:	Parameter	Options
	Solar EUV flux	 EUVAC (empirical model based on observed F_{10.7}) TIMED/SEE 37-band observational product SDO/EVE 37-band observational product
	Thermospheric composition	 MSIS empirical model (all versions) TIEGCM physics-based model output files COMPASS composition assimilation module
	Neutral winds	 HWM empirical model (all versions) TIEGCM physics-based model output files EMPIRE wind estimation module
	Initial electron density	 IRI empirical model (all versions) TIEGCM physics-based model output files Ionospheric assimilation module Previous output
	Ionospheric temperatures	IRI empirical model (all versions)TIEGCM physics-based model output files
	Ion drifts	 Scherliess-Fejer equatorial ionospheric drift model

Continuity equation (solved at 1 Hz):

production + loss + **∇**. (parallel flux + perpendicular flux)

Production rate (recalculated every five minutes):

$$q = \sum_{s} N_{s} \sum_{w} \sigma^{sw} \phi^{w} exp(-\tau^{w})$$

- where s: neutral species (normally O, O_2 , N_2)
 - N: number density
 - w: EUV wavelength band (37 EUVAC-style bands)
 - σ : ionization cross section
 - ϕ : Solar flux
 - τ : Optical depth (using Chapman grazing angle approximation)

Continuity equation (solved at 1 Hz):

production + **loss** + **∇**. (parallel flux + perpendicular flux)

Kirchengast [1996] loss rate (recalculated every second):

$$\frac{\beta_1 + \beta_2}{1 + \frac{\beta_1}{\alpha_1 N_e} + \frac{\beta_2}{\alpha_3 N_e}} N_e - \frac{q_{N_2}}{\beta_3 + \alpha_2 N_e} \frac{1 + \beta_3}{\alpha_1 N_e} - \frac{q_{O_2}}{\alpha_3 N_e}$$

where $\alpha_1 = 4.2\text{E}-13 \times (300 / \text{T}_e)^{0.85}$ $\alpha_2 = 1.8\text{E}-13 \times (300 / \text{T}_e)^{0.39}$ $\alpha_3 = 4.2\text{E}-13 \times (300 / \text{T}_e)^{0.55}$ $\beta_1 = N_{N_2} y_1$ $\beta_2 = N_{O_2} y_2$ $\beta_3 = N_{O} y_3$ $y_1 = 1.533\text{E}-12 - 5.920\text{E}-13 \times \text{T}_i / 300 + 8.600\text{E}-14 \times (\text{T}_i / 300)^2 \text{ below 1700 K}$ $= 2.730\text{e}-12 - 1.155\text{e}-12 \times \text{T}_i / 300 + 1.483\text{E}-13 \times (\text{T}_i / 300)^2 \text{ above 1700 K}$ $y_2 = 2.82\text{e}-11 - 7.74\text{e}-12 \times \text{T}_i / 300 + 1.073\text{e}-12 \times (\text{T}_i / 300)^2 - 5.17\text{e}-14 \times (\text{T}_i / 300)^3 + 9.65\text{e}-16 \times (\text{T}_i / 300)^4$ $y_3 = 1.1410\text{e}-10 \times (300 / \text{T}_i)^{0.44}$

Continuity equation (solved at 1 Hz):

production + loss + ∇. (parallel flux + perpendicular flux)

Parallel flux is composed of the field-aligned wind, gravity and diffusion:

$$N_{e} \cdot \left\{ u_{||} + \frac{g_{||}}{v_{in}} + \frac{k_{B}(T_{i} + T_{e})}{m_{i}v_{in}} \frac{\nabla_{||}N_{e}}{N_{e}} \right\}$$

where v_{in} : ion-neutral collision frequency

- $u_{||}$: magnetic meridional wind
- $g_{||}$: Field-aligned component of gravity
- k_B : Boltzmann's constant
- T: Temperature
- *m*: mass
- 7: Gradient calculated using a spline derivative approach

The diffusion term is simplified by neglecting thermal diffusion coefficients and the fieldaligned component of the ion stress tensor.

Continuity equation (solved at 1 Hz):

production + loss - ∇ · (parallel flux + perpendicular flux)

Perpendicular flux is calculated from specified equatorial ion drifts: $N_e. - \mathbf{v}_{E \times B} (cosl\hat{r} + sin |\hat{\theta})$

where \mathbf{v}_{ExB} : ion-neutral collision frequency *I*: inclination \widehat{r} : Radial unit vector $\widehat{\theta}$: Meridional unit vector

The perpendicular flux term is simplified by neglecting the zonal component. Velocities are set to zero poleward of 45°



Ionospheric assimilation module

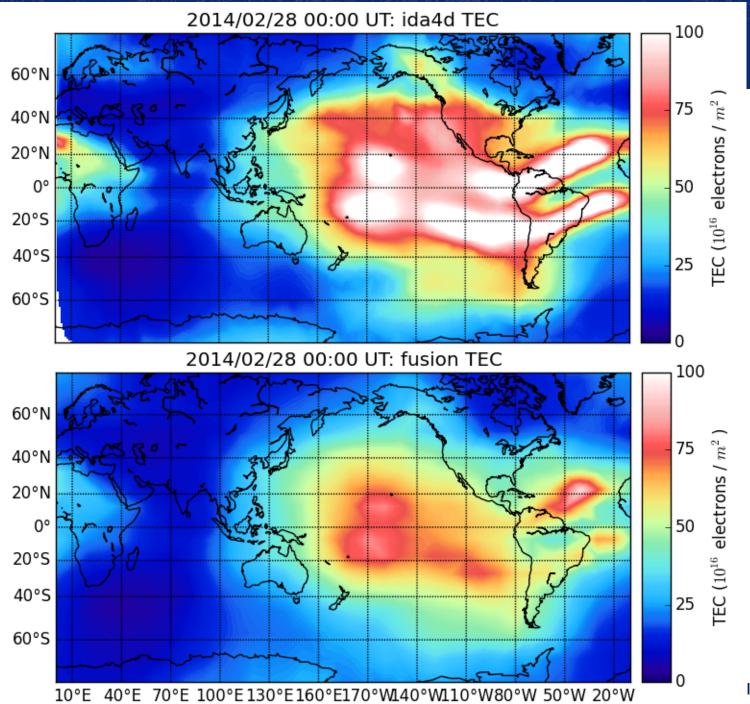
Ionospheric Data Assimilation Four-Dimensional (IDA4D) (*Bust and Datta-Barua*, [2014])

Assimilates ionospheric observations from ground- and space-borne GNSS receivers, HF instruments, Beacon transmitters, *in situ* densities, satellite radiances and other data sources.

Standalone mode, temporal advance is achieved using a Gauss-Markov Kalman Filter approach with IRI as an empirical background model

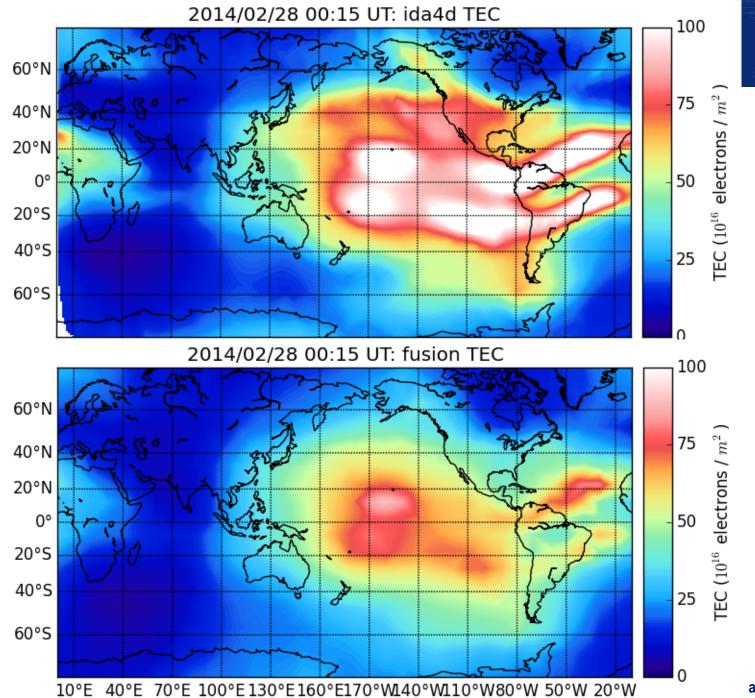
Coupled physics mode: temporal advance achieved using the physics module. Data are assimilated to the model background at each timestep

Bust, G. S., & Datta-Barua, S. (2014). Scientific investigations using IDA4D and EMPIRE. *Modeling the Ionosphere-Thermosphere System*, 283-297.



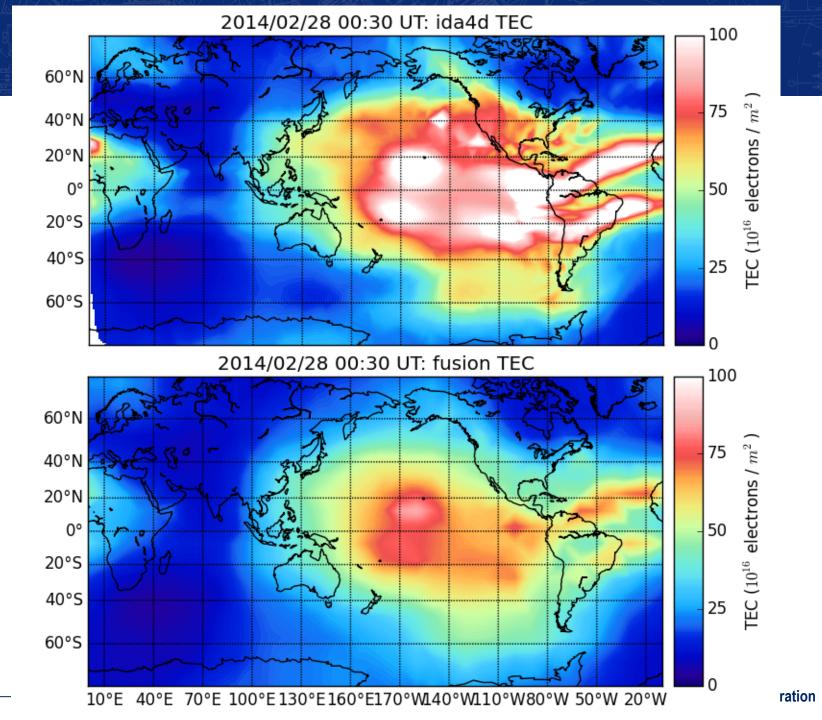
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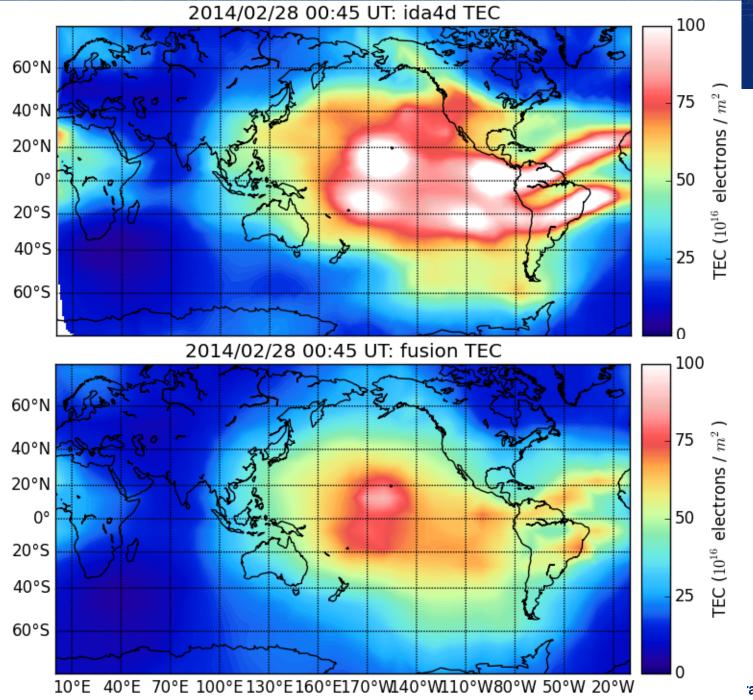


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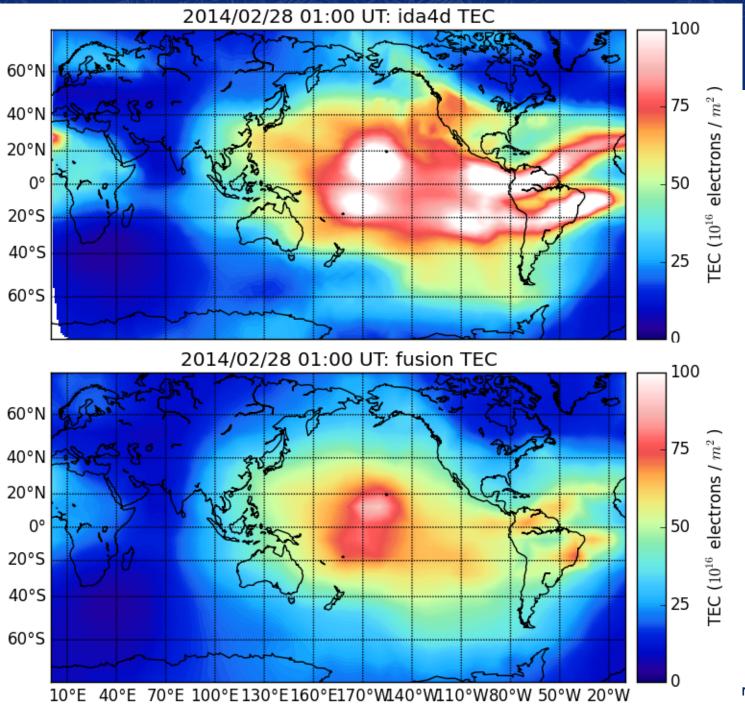


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Assimilates thermospheric O/N₂ observations from SSUSI and (soon) GUVI MSIS empirical background used in 3DVar approach O/N₂ is calculated in a column down to a depth that contains $10^{21} N_2 / m^2$

 $J = (x - x_b)^T B^{-1} (x - x_b) + (y - H x)^T R^{-1} (y - H x)$

- J cost function minimized using Powell's method
- $\mathbf{x}_{a} O/N_{2}$ ratio analysis for which J is minimum
- $\mathbf{x}_{b} O/N_{2}$ ratio background

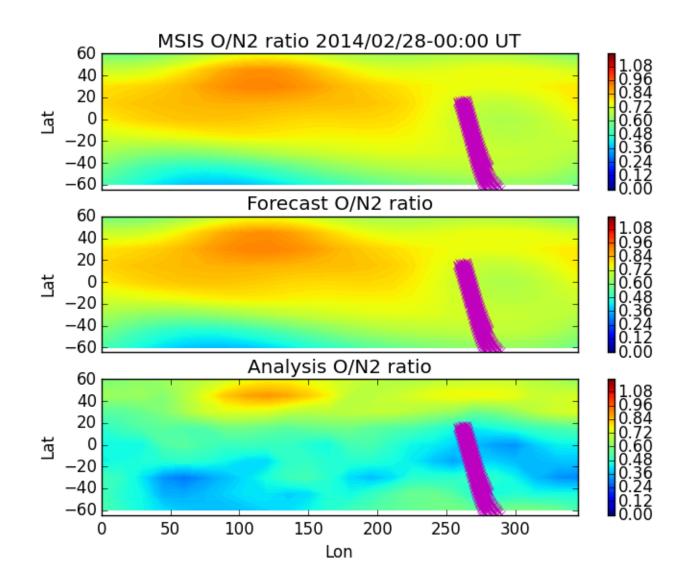
B – Background error covariance matrix calculated from a 10-year MSIS run (8x daily, 50 days between 2000 and 2010)

R – The (diagonal) observation error covariance as reported by SSUSI/GUVI

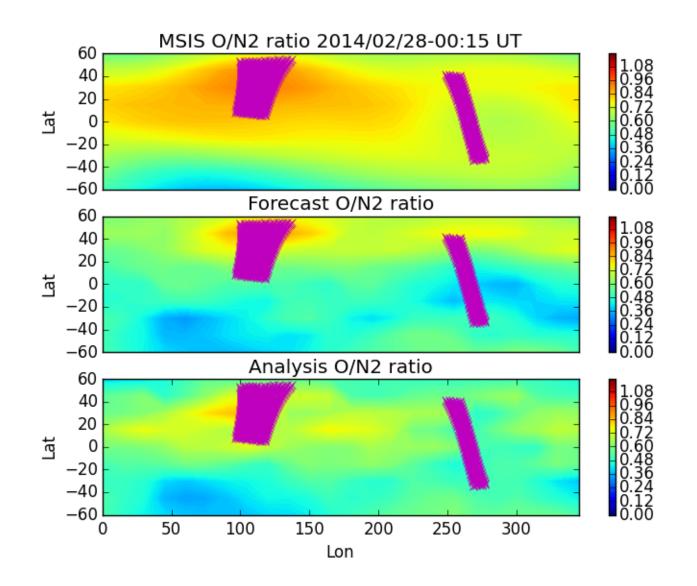
H - A linear observation operator that interpolates x_b to the observation locations using inverse-distance weighting of its four nearest neighbours

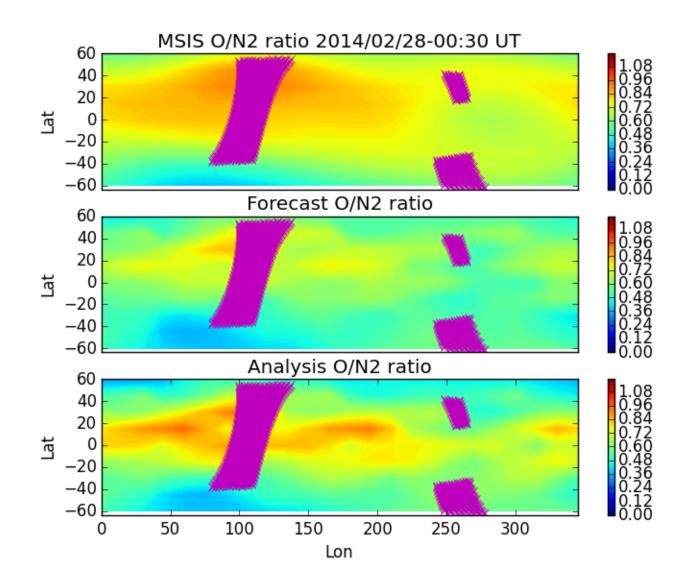
 \mathbf{x}_{b} is propagated forwards in time as follows:

 $\mathbf{x}_{b t2} = 0.9 (\mathbf{x}_{b t1} + MSIS_{t2} - MSIS_{t1}) + 0.1 MSIS_{t2}$

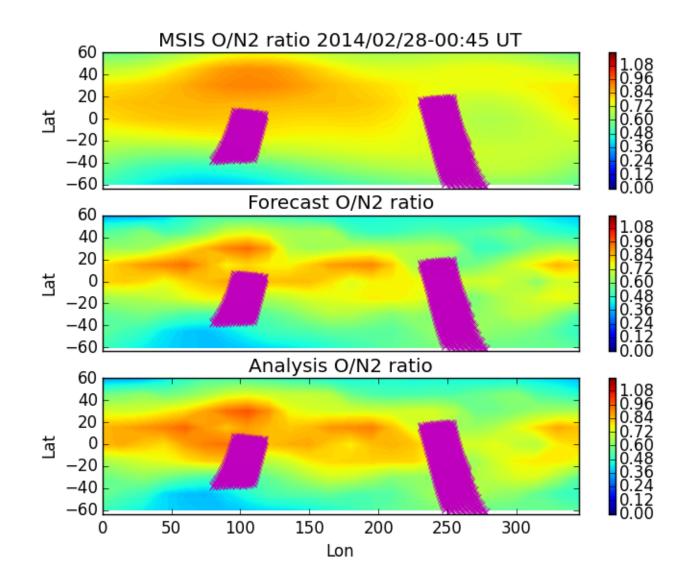


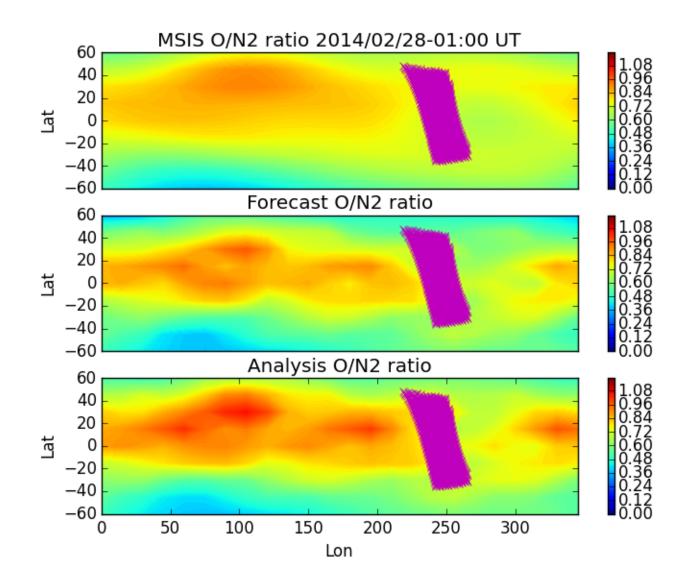
Space Exploration APL

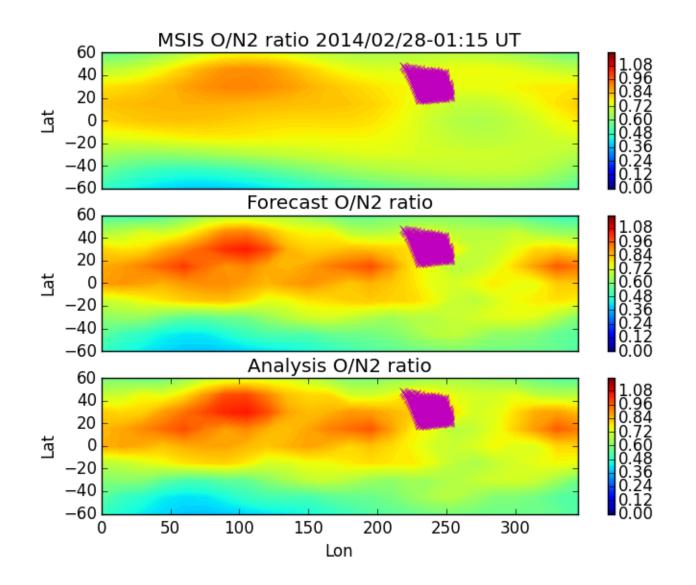




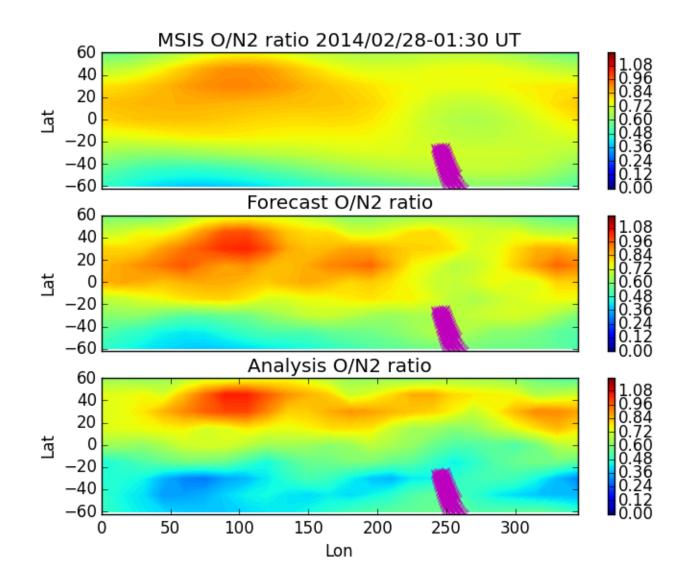




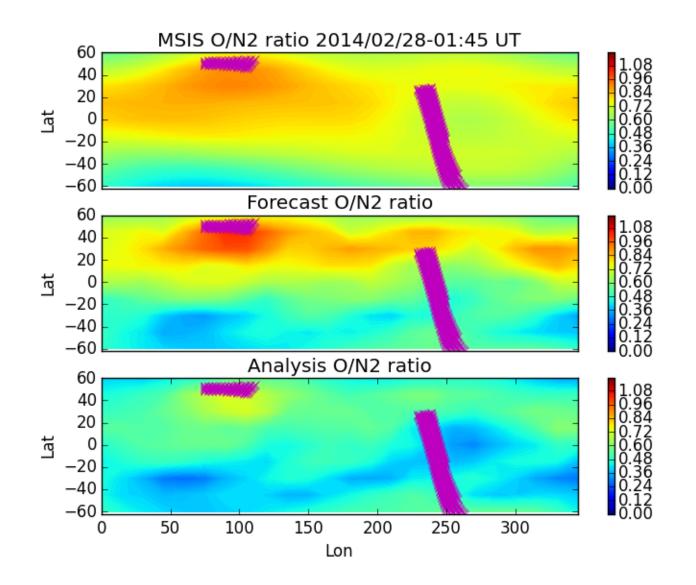




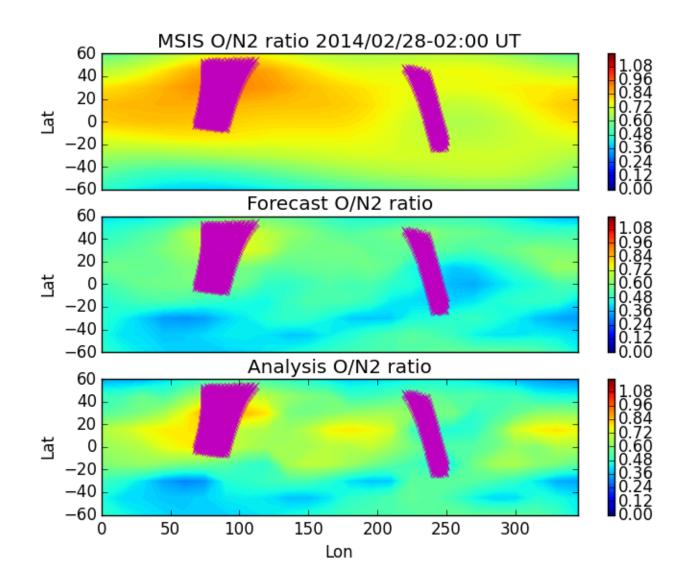
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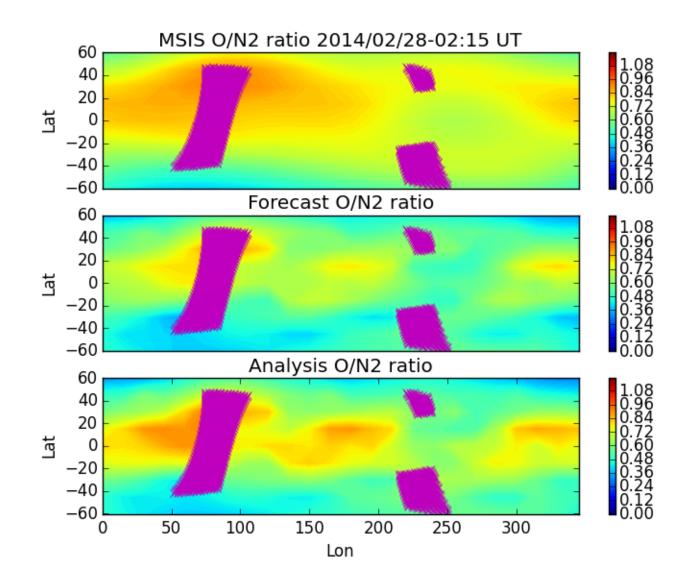
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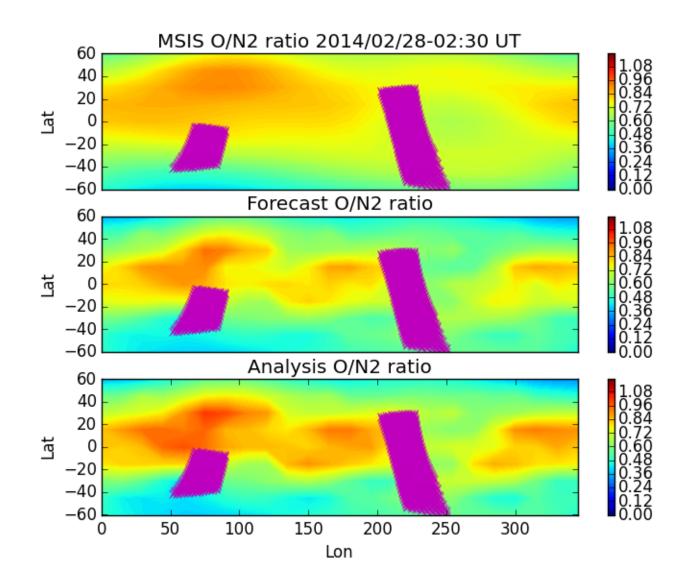


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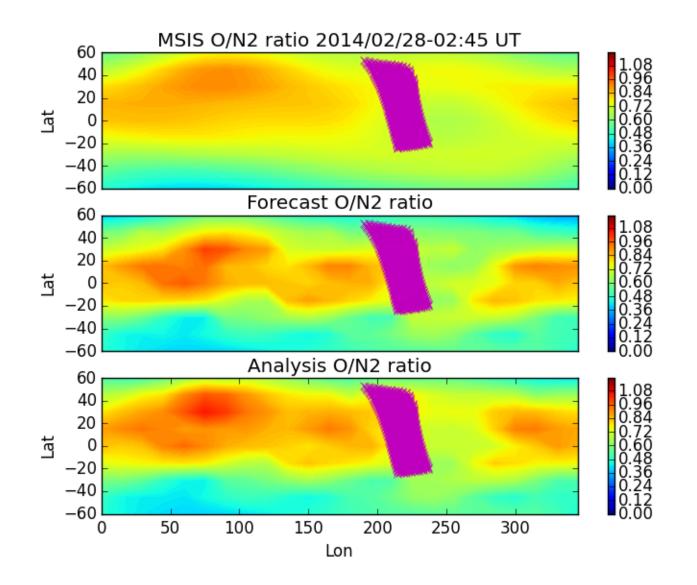


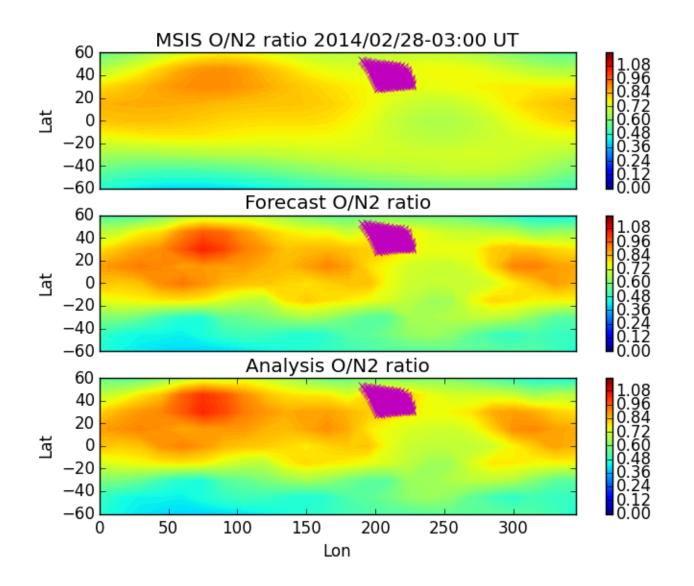
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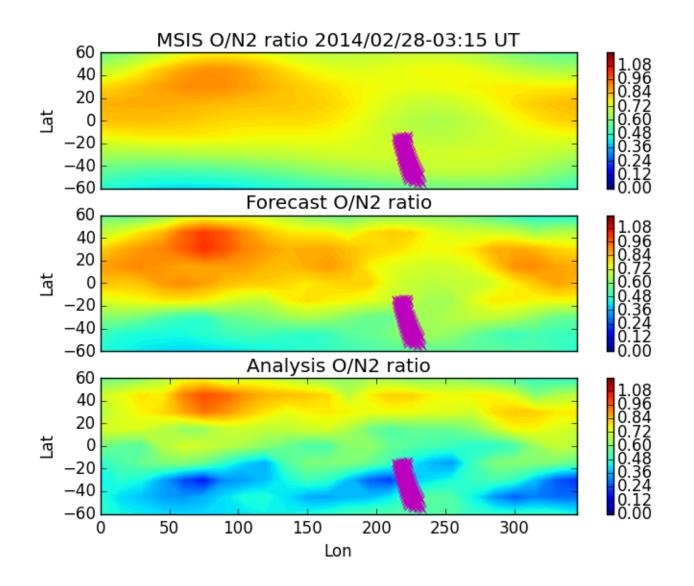




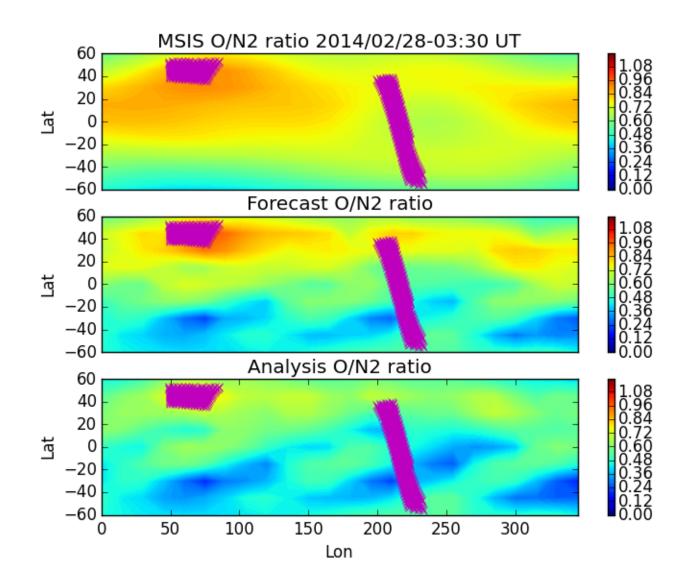
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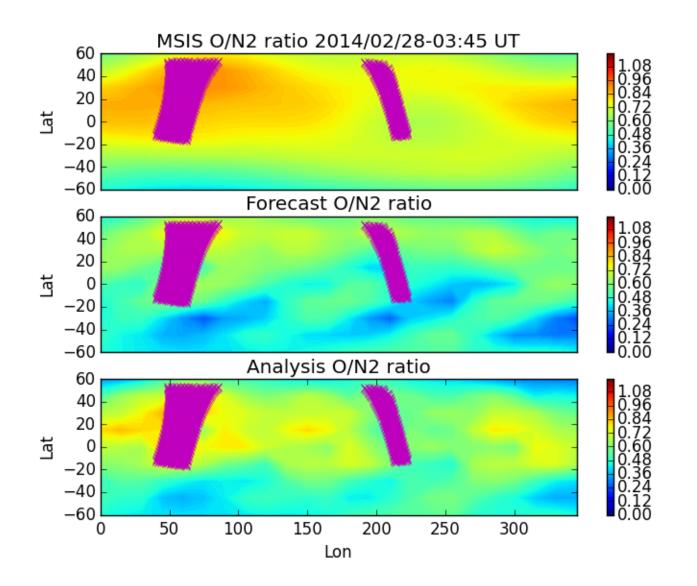


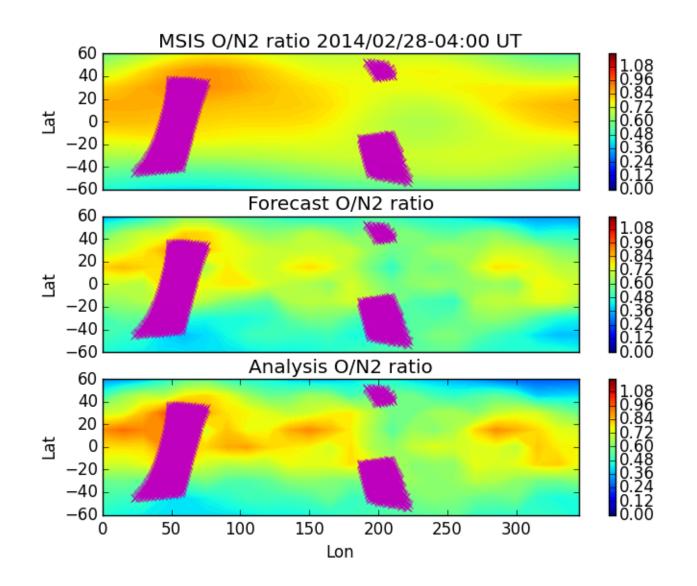




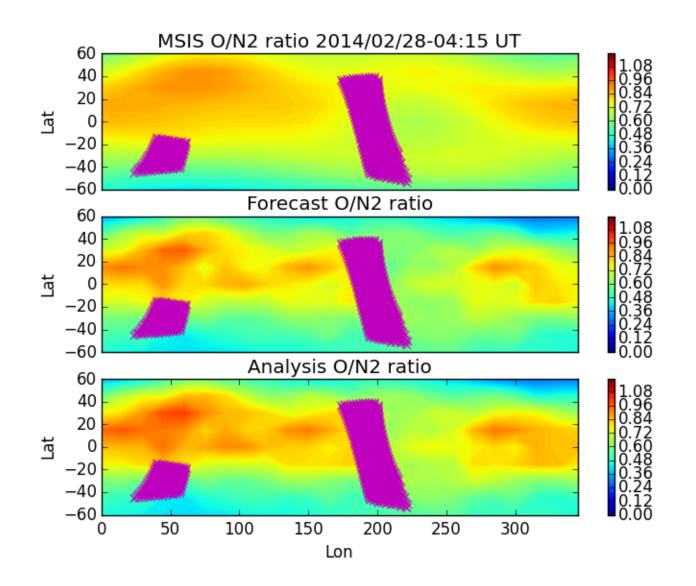
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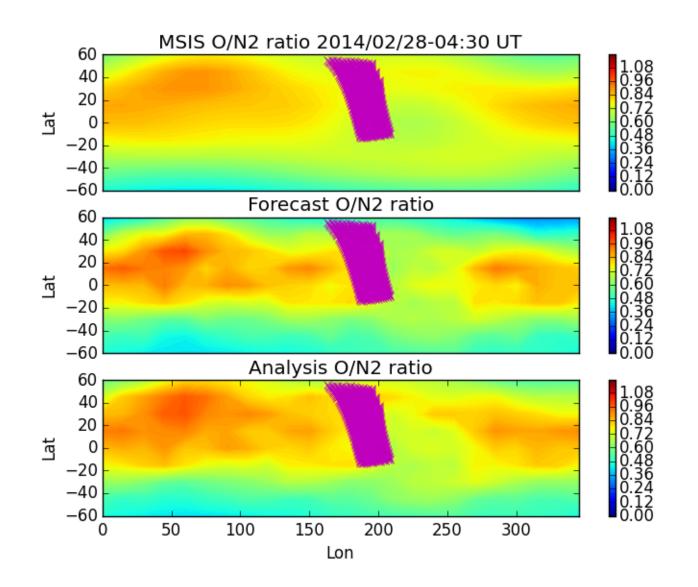


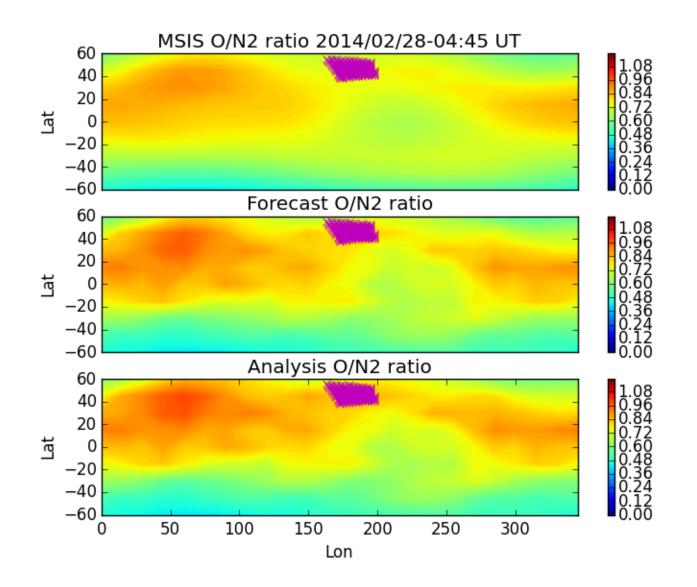


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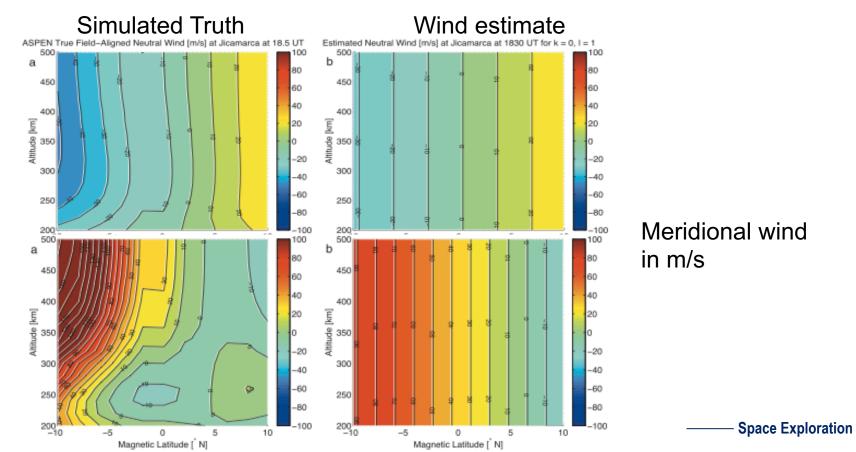
Space Exploration APL





Wind estimation module

- Estimating Model Parameters through lonospheric Reverse Engineering (EMPIRE) Datta-Barua et al. [2009]
- Meridional wind correction determined to explain the discrepancy between the electron density background and analysis
- All model terms are calculated using ionospheric physics described earlier

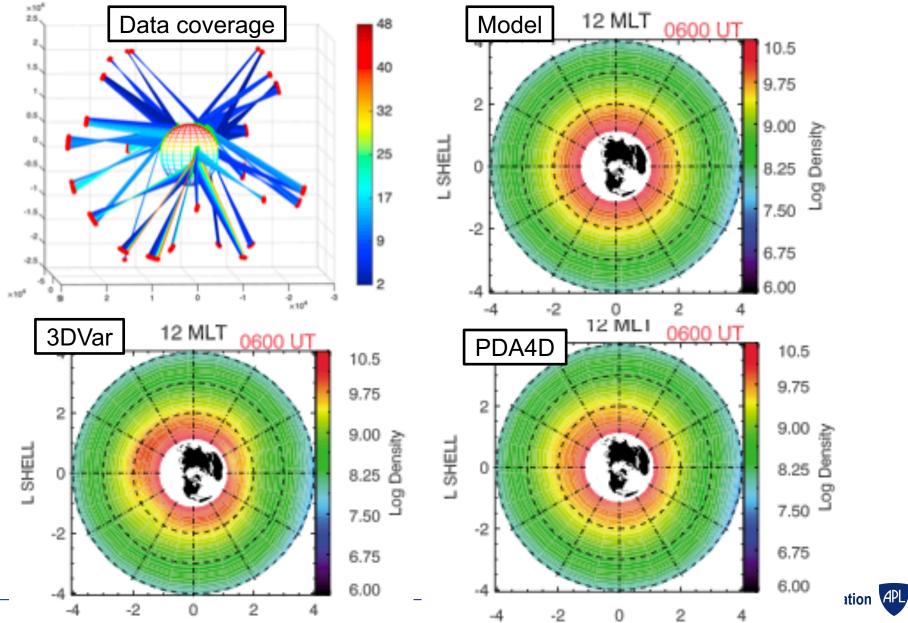


- Plasmaspheric Data Assimilation Four-Dimensional (*Nikoukar et al.* 2015)
- GPS data assimilated to Global Core Plasmasphere model
- Gauss-Markov Kalman Filter approach (3DVar also tested)

Nikoukar, R., Bust, G., & Murr, D. (2015). A novel data assimilation technique for the plasmasphere. *Journal of Geophysical Research: Space Physics*, *120*(10), 8470-8485.



Plasmaspheric assimilation module





- IDA2017 is a next-generation coupled, modular assimilation package
- Diverse modeling and data assimilation tools available and interchangeable through a common interface
- New developments include a physics advance module and a composition assimilation module early results shown

